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# NONLINEAR BREAKING-WAVE CHARACTERISTICS IN LAGRANGIAN COORDINATES

Wen-Jer Tseng<sup>1</sup> and Chia-Yan Cheng<sup>2</sup>

Key words: Eulerian system, Lagrangian system, breaking-wave characteristics.

### ABSTRACT

This paper presents a theoretical investigation of nonlinear surface-wave propagation over a sloping bottom. For a problem with nonlinear surface-wave propagation over a sloping bottom, a perturbation method is first used to find the analytical solution in order to derive the third order terms for the bottom slope  $\alpha$  and the wave steepness  $\varepsilon$  in the Eulerian system. Then, by transforming the flow field solution from the Eulerian system into the Lagrangian system, a more accurate wave profile is identified. By using the kinematic stability parameter, new theoretical breaking-wave characteristics are derived. The theoretical solutions are then compared with those from other research. The results reveal that the present solution reasonably describes the wave-breaking phenomenon. In this paper, a new theoretical solution for the breaking-wave characteristics is provided, and it is a useful approach for predicting breaking-wave characteristics.

### I. INTRODUCTION

Because of changes in water depth, a wave touches the bottom when the depth is shorter than half of its wavelength, and the wave shoals when travelling from deep water to shallow water. The wave height increases, and the wave profile becomes more skewed and asymmetric (Elgar and Guza, 1985) owing to the nonlinear effects. The celerity is reduced; hence, the fluid particle velocity of the wave crest is faster than the celerity, and the wave breaks. A breaking wave releases enormous energy and substantial impact force, which damages coastal structures such as breakwaters and causes sediment transport along and across nearby shores. Therefore, break-

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ing-wave characteristics must be quantified to obtain reliable predictions about sediment transport and structural design in coastal regions. Most investigators have discussed and inducted formulae from experimental studies. For example, Street and Camfield (1966), Tang (1966), L'e M'ehaut'e and Koh (1967), Galvin (1968), Goda (1970, 2004), Saeki et al. (1971), Sunamura and Horikawa (1974), Sunamura (1980, 1983), Ogwa and Shuto (1984), Seyama and Kimura (1988), Hansen (1990), and Rattanapitikon and Shibayama (2000) have all proposed empirical formulae or semiempirical formulae for breaking-wave height and breaking-wave depth as a function of deep-water wave height, wave steepness, and bottom slope, in accordance with experiments or reliable field data. Chang (1999), Gotoh and Sakai (1999), and Hsieh et al. (2007, 2008) have studied breaking-wave characteristics by means of numerical simulation, and Deo et al. (2003) studied the same topic by means of neural networks and experimental results. Iwagaki et al. (1974), Chanson and Lee (1997), Ting et al. (2002), Tsai et al. (2005), and Hsiao et al. (2008) have used experimental data to examine some empirical formulae and characteristics related to breaking waves. The major drawback for all of the aforementioned studies has been that they could only derive the breaking-wave height or breaking-wave depth and the location of a wave when breaking; however, these studies have not been able to completely describe the process of the wave deformation and of the related flow field.

Biesel (1952) suggested a plausible approximation method to account for normal incident waves propagating on a sloping bottom, with the bottom slope taken into account in the velocity potential as a perturbation parameter, but wave breaking and nonlinear effects are not addressed in that study. Furthermore, Biesel did not explain the derivation of the formula in detail. Keller (1958), Chen and Tang (1992), Hsu et al. (2001), and Chen et al. (2005, 2006) have extended Biesel's theory.

In this study, a nonlinear analysis was undertaken by perturbing to the third order of the bottom slope  $\alpha$  or the wave steepness  $\varepsilon$  in an Eulerian system. Then, the profile of the breaking wave was determined using a transformation into a Lagrangian system. The breaking-wave height and the breaking wave depth can be derived using kinematic stability

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Fig. l. Sketch definition for a surface wave propagating on a gently sloping bottom.

parameters. Finally, the theoretical solutions were verified by comparing them with those from other studies, as shown in Fig. 3 to Fig. 5.

# **II. MATHEMATICAL FORMULATION**

To describe a surface wave propagating toward a gently and uniformly sloping bottom, a two-dimensional Cartesian x - ycoordinate system is used, as shown in Fig. 1. The negative *x*-axis is directed outward to the sea, and the positive *y*-axis extends vertically upward from the still water level; the bottom, or the seabed, is at  $y = -h = \alpha x$ , where  $\alpha$  denotes the bottom slope (Li et al., 2013).

In Fig. 1, angle  $\beta$  is the incline of the bottom,  $\alpha (= \tan \beta)$  is the bottom slope,  $\eta(x,t)$  is the elevation of the water's surface, C is celerity, h(x) is water depth,  $H_b$  is breaking-wave height, and  $h_b$  is breaking-wave depth.

Assume that the flow field is irrotational, incompressible, and inviscid. The governing equations and boundary conditions can be derived to the third order of the bottom slope  $\alpha$ or of the wave steepness  $\varepsilon (= H_o / L_o)$ , where  $H_o$  is the wave height and  $L_o$  is the wavelength of a deep-water wave or an incident wave. To describe the irrotational motion of an inviscid and incompressible fluid, a velocity potential  $\phi(x, y, t)$  is introduced, and the horizontal and vertical velocities are given by

$$u = \frac{\partial \phi}{\partial x}, \ v = \frac{\partial \phi}{\partial y}, \ \vec{V} = u\vec{i} + v\vec{j}.$$
(1)

The velocity potential  $\phi(x, y, t)$  is harmonic with respect to x and y and satisfies the continuity equation. This leads to the Laplace equation

$$\nabla^2 \phi = \frac{\partial^2 \phi}{\partial x^2} + \frac{\partial^2 \phi}{\partial y^2} = 0$$
 (2)

The wave motion described must satisfy boundary conditions at the bottom and at the free surface, respectively, as follows:

1. On an immovable and impermeable sloping bottom with

an incline  $\beta$  to the horizon, the bathymetry is represented by  $f(x, y) = y + h = y - \alpha x = 0$ , and the no-flux bottom boundary condition gives

$$\nabla \phi \cdot \overline{n} = \nabla \phi \cdot \frac{\nabla f}{|\nabla f|} = \frac{-\alpha \phi_x + \phi_y}{\sqrt{1 + \alpha^2}} = 0, \ y = -h, \ \alpha > 0 \text{ and } x \le 0,$$

or simply

$$\frac{\partial \phi}{\partial y} - \alpha \frac{\partial \phi}{\partial x} = 0, \ y = -h \ . \tag{3}$$

2. The dynamic and kinematic free-surface boundary conditions are

$$\phi_t + (\nabla \phi)^2 / 2 + g\eta = 0, \ y = \eta$$
 (4)

$$\phi_{y} = d\eta / dt, \ y = \eta \tag{5}$$

Taking the total differential D/Dt of Equation (4) and using Equation (5) gives

$$\phi_{tt} + g\phi_{y} + [(\nabla\phi)^{2}]_{t} + \nabla\phi \cdot \nabla(\nabla\phi)^{2} / 2 = 0, y = \eta$$
(6)

3. A time-averaged mass flux conservation condition is required: The water mass bounded by the shoreline created by the sloping bottom rushes up and down the beach face, and the total mass does not change. Therefore, at any cross-section where x is a constant, the time-averaged mass flux should vanish. This condition gives

$$y - direction: \frac{1}{T} \int_0^T \int_{-h}^{h} \phi_y \, dy dt =$$

$$\overline{\int_{-h}^{h} \phi_y \, dy} = \overline{(\phi)_{y=\eta}} - \overline{(\phi)_{y=-h}} = 0,$$
(7)

$$x - direction: \frac{1}{T} \int_{0}^{T} \int_{-h}^{\eta} \phi_x \, dy dt - \frac{U(\alpha)}{T} \int_{0}^{T} \int_{-h_0}^{\eta^c} \phi_x^c \, dy dt$$
$$= \overline{\int_{-h}^{\eta} u \, dy} - U(\alpha) \overline{\int_{-h_0}^{\eta^c} u^c \, dy} = 0; U(\alpha) = \begin{cases} 0, \alpha \neq 0\\ 1, \alpha = 0 \end{cases}$$
(8)

Both the superscript c and the sub script 0 express the physical quantity at infinite depth. Because of the nonlinear effect, waves over constant depth induce a net flux of water. In Equation (8), a deep-water streaming term is introduced that is adjusted by a unit function  $U(\alpha)$  to ensure that it can be reduced to the deep-water condition when the bottom slope is equal to zero.

For a gentle bottom slope  $\alpha$ , Equations (2) to (8) can be solved order by order in the following section.

# **III. ASYMPTOTIC SOLUTIONS**

In this section, an explicit expression for the velocity potential of the wave field is first derived as a function of the bottom slope  $\alpha$  and of the wave steepness  $\mathcal{E}$  to the third order in an Eulerian coordinate system. Then, the wave profile and the velocity components are transformed into a Lagrangian system. Finally, in the next section, the kinematic stability parameter is introduced and theoretical breaking-wave characteristics derived. The detailed computation is as follows:

First, the Laplace equation, Equation (2), and the corresponding boundary conditions, Equation (3) to (8), are used to obtain the asymptotic solutions. This study assumes that the relevant physical quantities can be expanded as a double power series of the parameter  $\varepsilon$  and the bottom slope  $\alpha$ . Thus, the velocity potential  $\phi$ , the free-surface displacement  $\eta$ , the wave number k, and the angular frequency  $\sigma$  can be obtained as follows:

$$\sigma = \sum_{\beta=0}^{\infty} \varepsilon^{\beta} \sigma_{\beta}(\lambda_0, a_0) = \text{constant}$$
(9)

$$k = \sum_{r=0}^{\infty} \sum_{l=0}^{\infty} \varepsilon^r \alpha^l k_{r,l}(k_{0,0}, k_{0,0}h; \lambda_0, a_0)$$
(10)

$$\eta = \sum_{m=1}^{\infty} \varepsilon^{m} \eta_{m} (\lambda h, S; \alpha, \lambda_{0}, a_{0})$$

$$= \sum_{m=1}^{\infty} \sum_{n=0}^{\infty} \varepsilon^{m} \alpha^{n} \eta_{m,n} (\lambda h, S; \lambda_{0}, a_{0})$$

$$= \sum_{m=1}^{\infty} \sum_{n=0}^{\infty} \sum_{i=0}^{m} \varepsilon^{m} \alpha^{n} \eta_{m,n,i} (\lambda h, S; \lambda_{0}, a_{0})$$
(11)

$$\begin{split} \phi &= \sum_{m=1}^{\infty} \sum_{n=0}^{\infty} \varepsilon^m \alpha^n (\phi'_{m,n} + \phi_{m,n}) \\ &= \sum_{m=1}^{\infty} \sum_{n=0}^{\infty} \varepsilon^m \alpha^n \{\phi'_{m,n}(t, \sigma_0, a_0) \\ &+ \sum_{i=0}^m A_{m,n,i}(\lambda h, \lambda y, \lambda; \sigma_0, a_0) F_{m,n,i}(S) \times E_m(\lambda h, \lambda; \alpha) \\ &+ \int_{x_0}^x M_{m,n,0}(\lambda h, \lambda; \sigma_0, a_0) dx' \} \\ S &= \int_{x_0}^x k dx' - \sigma t ; E_m(\lambda h, \lambda, \alpha) \\ &= \exp[m \sum_{j=2}^{\infty} \alpha^j \int_{x_0}^x e_j(\lambda h, \lambda) dx'] \\ &= E_m; \lambda = k, \end{split}$$
(12)

where S is the phase function,  $x_0$  denotes the X value at infinite depth,  $e_j$  is the wave amplitude influence factor, and  $M_{m,n,0}$  is the return flow. The bottom slope  $\alpha$  is assumed to be small; thus, the qth differentiations of  $k_{r,l}$ ,  $A_{m,n,i}$ , and  $M_{m,n,0}$  with respect to X are of order  $\alpha^q$ :

$$\left(\frac{d^{q}k_{r,l}}{dx^{q}}, \frac{d^{q}e_{j}}{dx^{q}}, \frac{\partial^{q}A_{m,n,i}}{\partial x^{q}}, \frac{d^{q}M_{m,n,0}}{dx^{q}}\right) = O(\alpha^{q})$$
(13)

Substituting Equations (9) to (13) into the governing equation and the boundary conditions (2) to (8), using the Taylor series expansion of functions at y = 0 instead of  $y = \eta$ , and collecting terms of the same order of  $\varepsilon$  and  $\alpha$  yields the required equations to each order of approximation. Then different orders of  $\varepsilon(m)$ ,  $\alpha(n)$ , and harmonic (*i*) may be separated, yielding a set of partial differential equations for each index (m, n, i). The boundary-value problems can be solved sequentially for orders of m, n, and i. The solution is as follows, where the notation shown is used for convenience:

$$\cosh = \cosh, \sinh = \sinh, \tanh = \th; I = \coth \lambda h$$
 (14)

The solutions are obtained as follows:

For  $O(\varepsilon^2 \alpha)$ , the solution for velocity potential function  $\phi$  and the corresponding free-surface displacement  $\eta$  can be obtained from the boundary-value problem with corresponding order terms. The results are

$$\phi = \sum_{m=1}^{\infty} \sum_{n=0}^{\infty} \varepsilon^{m} \alpha^{n} (\phi'_{m,n} + \phi_{m,n})$$
  
=  $[(A_{1,0,1} + \alpha^{2} A_{1,2,1}) \sin S + (\alpha A_{1,1,1} + \alpha^{3} A_{1,3,1}) \cos S]$   
+  $[A_{2,0,2} \sin 2S + \alpha (A_{2,1,2} \cos 2S + A_{2,1,0})]$   
+  $[\phi'_{2,0} + \int_{x_{0}}^{x} M_{2,0,0} dx']$  (15)

$$\eta = \sum_{m=1}^{\infty} \sum_{n=0}^{\infty} \varepsilon^{m} \alpha^{n} \eta_{m,n}$$
  
= [(\alpha a\_{1,1,1} + \alpha^{3} a\_{1,3,1}) \sin S + (a + \alpha^{2} a\_{1,2,1}) \cos S] (16)  
+ [\alpha a\_{2,1,2} \sin 2S + a\_{2,0,2} \cos 2S] + \eta\_{2,0,0}

$$\sigma = \sigma_0; \, k = k_0 + \alpha^2 k_{0,2}, \tag{17}$$

where

$$A_{1,0,1} = \frac{ag}{\sigma_0} \frac{\mathrm{ch}\lambda(h+y)}{\mathrm{ch}\lambda h};$$
$$a_{1,1,1} = a(\frac{\lambda^2 h^2}{D\mathrm{sh}2\lambda h} - \lambda h + \frac{I^2}{D^2}\lambda h)$$

$$\begin{split} A_{1,1,1} &= \frac{ag}{\sigma_0} \left\{ \left[ -\frac{\lambda^2 (h+y)^2}{D \operatorname{sh} 2\lambda h} + \lambda (h+y) \right] \right. \\ &\times \frac{\operatorname{ch} \lambda (h+y)}{\operatorname{ch} \lambda h} - \frac{\lambda (h+y)}{D^2 \operatorname{th} \lambda h} \frac{\operatorname{sh} \lambda (h+y)}{\operatorname{ch} \lambda h} \right\} \\ A_{1,2,1} &= \frac{ag}{\sigma_0} \left\{ \left[ -\frac{\lambda^4 (h+y)^4}{2 (D \operatorname{sh} 2\lambda h)^2} + \frac{\lambda^3 (h+y)^3}{D \operatorname{sh} 2\lambda h} \right. \\ &\left. - \frac{(D^4 - 4D + 5I^2)}{2D^4} \lambda^2 (h+y)^2 + \frac{I\lambda (h+y)}{D^2} \right] \right] \\ &\times \frac{\operatorname{ch} \lambda (h+y)}{\operatorname{ch} \lambda h} + \left[ -\frac{4D + 10 \operatorname{ch}^2 \lambda h}{3D^3 \operatorname{sh}^2 2\lambda h} \lambda^3 (h+y)^3 \right. \\ &\left. + \frac{2(D + \operatorname{ch}^2 \lambda h)}{D^2 \operatorname{sh} 2\lambda h} \lambda^2 (h+y)^2 - (\lambda - k_{0,2})(h+y) \right] \\ &\times \frac{\operatorname{sh} \lambda (h+y)}{\operatorname{ch} \lambda h} \right\} \end{split}$$

$$a_{1,2,1} = a \{ -\frac{(\lambda h)^4}{2(D \mathrm{sh} 2\lambda h)^2} + \frac{(\lambda h)^3}{D \mathrm{sh} 2\lambda h} -\frac{(4D + 10 \mathrm{ch}^2 \lambda h)(\lambda h)^3}{3D^3 I \mathrm{sh}^2 2\lambda h} + \left[ -\frac{D^4 - 4D + 5I^2}{2D^4} + \frac{D + \mathrm{ch}^2 \lambda h}{D^2 \mathrm{ch}^2 \lambda h} \right] (\lambda h)^2 + \left[ \frac{I}{D^2} - (1 - \frac{k_{0,2}}{\lambda}) / I \right] \lambda h \}$$

$$k_{0,2} = [D^{5} + (D^{4} - 18D^{2} + 32D - 15)ch^{2}\lambda h + 2D^{2} + (-3D^{2} + 15D - 15)I^{2}]\lambda/(3D^{5})$$

 $a = a_0 / \sqrt{D \operatorname{th} \lambda h}$ ;  $D = 1 + 2\lambda h / \operatorname{sh} 2\lambda h$ 

$$A_{2,0,2} = \left[\frac{3}{8}\sigma_0 a^2 \frac{\operatorname{ch} 2\lambda(h+y)}{\operatorname{sh}^4 \lambda h}\right]; \ \phi'_{2,0} = -\sigma_0^2 a_0^2 t / (4\operatorname{sh}^2 \lambda_0 h_0)$$
$$M_{2,0,0} = \left[\frac{-g\lambda^2 a^2}{2\sigma_0 \lambda h} + U(\alpha)g\lambda\lambda_0 a_0^2 / (2\sigma_0\lambda h)\right]; \ \eta_{2,0,0} = \overline{\eta_{2,0}}$$
$$= (\lambda_0 a_0^2 / \operatorname{sh} 2\lambda_0 h_0 - \lambda a^2 / \operatorname{sh} 2\lambda h) / 2 \le 0$$

$$a_{2,0,2} = \frac{1}{4} \lambda a^2 \frac{\operatorname{ch} \lambda h}{\operatorname{sh}^3 \lambda h} (2 \operatorname{sh}^2 \lambda h + 3)$$

$$\begin{aligned} a_{2,1,2} &= \lambda a^2 \{ (3I^3 - I)\lambda^2 h^2 / (2D \text{sh} 2\lambda h) + [12DI^4 \\ &+ (-3D^2 + 6D + 6)I^3 - 24DI^2 + (D^2 - 5D - 5)I \\ &+ 12D + (-D + 1)I^{-1} ]\lambda h / (2D^2) + [12D^2I^7 \\ &+ (6D^2 + 3D)I^6 + (-36D^2 + 24D)I^5 + (-12D^2 + D + 22)I^4 \\ &- 12D^2I^3 + (-18D^2 - 3D + 4)I^2 + (36D^2 - 24D)I \\ &+ (-8D^2 - D + 6)] / (8D^2) \end{aligned}$$

$$\begin{split} A_{2,1,2} &= \sigma_0 a^2 \{ [-\frac{3(I^4-1)\lambda^2(h+y)^2}{4D \operatorname{sh} 2\lambda h} + \frac{3(I^2+1)\lambda(h+y)}{4\operatorname{sh}^2\lambda h} \\ &- (I^2+1)(12D^2I^6 + (6D^2+3D)I^5 - (48D^2-24D)I^4 \\ &- (18D^2+2D-22)I^3 + (36D^2-24D)I^2 - (D+10)I) \\ &/ (16D^2)] \frac{\operatorname{ch} 2\lambda(h+y)}{\operatorname{ch} 2\lambda h} \\ &- \frac{3[2D(I^4-1) + (D+1)(I^3+I)]}{4D^2\operatorname{sh}^2\lambda h} \lambda(h+y) \\ &\times \frac{\operatorname{sh} 2\lambda(h+y)}{\operatorname{ch} 2\lambda h} \} \\ A_{2,1,0} &= \{ \frac{\sigma_0}{4D} [\frac{Ia^2}{(\lambda h)^2} + \frac{2}{\lambda h} (\frac{D+1}{D}I^2 - 2)a^2] - U(\alpha) \frac{g\lambda_0 a_0^2}{4\sigma_0 (\lambda h)^2} \} \\ &\times \lambda^2(h+y)^2 + \{ \frac{\sigma_0}{2D} [\frac{D-1}{\lambda h}I + \frac{1-2D}{D} \times (I^2-1)]a^2 \frac{e^{\alpha\lambda(h+y)}-1}{e^{\alpha\lambda h}-1} \\ &- \sigma_0 [\frac{Ia^2}{2\lambda h} - U(\alpha) \frac{\lambda_0 a_0^2 I}{2\lambda^2 h}] \} \times \lambda(h+y) + \{ \frac{\sigma_0}{4D} [(\frac{4D-2}{D}I^2 - 2)a^2 ] \\ &+ (1+\frac{2}{D}-2D) \times \frac{Ia^2}{\lambda h} ](\lambda h) - U(\alpha) \frac{g\lambda_0 a_0^2}{4\sigma_0} \} \frac{e^{\alpha\lambda(h+y)}-1}{e^{\alpha\lambda h}-1} \end{split}$$

$$\begin{split} A_{1,3,1} &= \frac{ag}{\sigma_0} \cdot \left\{ \frac{\lambda^6 (h+y)^6}{6D^3 \sinh^3 2\lambda h} - \frac{\lambda^5 (h+y)^5}{2D^2 \sinh^2 2\lambda h} + \left[ \frac{1}{2D \sinh 2\lambda h} \right] \right. \\ &+ \frac{D(12D + 40\cosh^2 \lambda h - 16D \sinh^2 2\lambda h + 100\cosh^4 \lambda h}{6D^5 \sinh^3 2\lambda h} \right] \cdot \lambda^4 (h+y)^4 \\ &- \left[ \frac{1}{6} + \frac{D(4D + 10\cosh^2 \lambda h) - 2D \sinh^2 2\lambda h + 10\cosh^4 \lambda h}{D^4 \sinh^2 2\lambda h} \right] \cdot \lambda^3 (h+y)^3 \\ &+ \left[ \frac{3D + (4 - 2k_{0,2}/\lambda) \cosh^2 \lambda h}{D^2 \sinh 2\lambda h} - \frac{k_{0,2x}}{2\alpha^4 \lambda^2} \right] \cdot \lambda^2 (h+y)^2 \\ &+ \left[ \frac{3D + (4 - 2k_{0,2}/\lambda) \cosh^2 \lambda h}{D^2 \sinh 2\lambda h} \cdot \lambda^5 (h+y) \right] \\ &+ \left[ \frac{4}{3D^2} \left\{ \frac{20D + 35\cosh^2 \lambda h}{15D^4 \sinh^3 2\lambda h} \cdot \lambda^5 (h+y)^5 \right\} \\ &- \left[ \frac{10D + 10\cosh^2 \lambda h}{D^2 \sinh 2\lambda h} \lambda^4 \cdot (h+y)^4 \right] \\ &+ \left[ \frac{3D + \cosh^2 \lambda h - k_{0,2}/\lambda \cdot D}{D^2 \sinh 2\lambda h} + \frac{4_{xxx}}{6\alpha^3} \right] \cdot \lambda^3 (h+y)^3 \\ &- \left[ \frac{D^4 - 4D + 5\cosh^2 \lambda h}{D^4} - \frac{k_{0,2}}{\lambda} \right] \cdot \lambda^2 (h+y)^2 \\ &+ \left[ \frac{2\cosh^2 \lambda h}{D^2 \sinh 2\lambda h} - e_3 \right] \cdot \lambda (h+y) \right\} \cdot \frac{\sinh \lambda (h+y)}{\cosh \lambda h} \end{split}$$

$$a_{1,3,1} = \frac{-\sigma}{g} \cdot \sum_{j=1} [M_{1,3,1,j} \cosh \lambda h + N_{1,3,1,j} \sinh \lambda h] \cdot h^j;$$
  
$$\sigma^2 = g\lambda \tanh \lambda h = g\lambda_0 \tanh \lambda_0 h.$$



Fig. 2. Comparison of wave profiles with Lagrangian and Eulerian coordinates.

$$a = \frac{a_0}{\left(D \tanh \lambda h\right)^{\frac{1}{2}}}; D = \left(1 + \frac{2\lambda h}{\sinh 2\lambda h}\right), \qquad (18)$$

In Equations (15) to (18),  $\sigma = \sigma_0 = \frac{2\pi}{T}$  is the angular frequency, *T* is the wave period, and  $a_0$  is the wave amplitude in deep water.

# **IV. BREAKING WAVES**

#### 1. Wave deformation

As a wave reaches its limit, the crest is fully developed as a summit and becomes highly asymmetrical. This asymmetric wave profile can be described by a Lagrangian coordinate system but cannot be captured by Eulerian solutions. Hence, the motion described in the Lagrangian system is close to that observed in reality (Chen et al., 2004).

The Eulerian and Lagrangian wave profile prior to breaking as a wave propagating over a gently sloping bottom  $(0 < \alpha \le 1/10)$  is illustrated in Fig. 2.

According to the velocity potential given in the previous section, the horizontal and vertical Eulerian velocity components,  $u_e = \partial \phi / \partial x$  and  $v_e = \partial \phi / \partial y$ , of a fluid particle can be derived. As a wave shoals, the wave form becomes asymmetric, and the particle motion can be obtained by the transformation between Eulerian and Lagrangian coordinates. The Lagrangian velocity may be estimated from the Eulerian velocity (Chen et al., 2006)  $\vec{V}_{L}(\vec{a},t) = \vec{V}_{E}[a\int_{t}^{t} \vec{V}_{L}(\vec{a},t') = \vec{V}_{E}(\vec{a},t) + \int_{t}^{t} \vec{V}_{L}(\vec{a},t')dt' \cdot \nabla_{\vec{a}}\vec{V}_{E}(\vec{a},t) + H.O.T.$ 

The Eulerian (Lagrangian) velocity is  $\vec{V}_E(\vec{a},t) = [u_e,v_e]$  $(\vec{V}_L(\vec{a},t) = [u_l,v_l])$ , and  $\vec{a}$  is the position of a fluid particle at time  $t = t_0$ . The Eulerian motions can then be transferred into the Lagrangian system according to the trajectory of the fluid particle. The corresponding velocity components  $u_l$  and  $v_l$  are

$$u_{l} = u_{e} + (\int^{t} u_{e} dt) \frac{\partial u_{e}}{\partial x} + (\int^{t} v_{e} dt) \frac{\partial u_{e}}{\partial y} = \sigma_{0} a [(\alpha C_{1} + \alpha^{3} C_{3}) \frac{ch\lambda(h+y)}{sh\lambda h} + (\alpha S_{1} + \alpha^{3} S_{3}) \times \frac{sh\lambda(h+y)}{sh\lambda h}] \sin S$$

$$+ \sigma_{0} a [(1 - \alpha^{2} C_{2}) \frac{ch\lambda(h+y)}{sh\lambda h} - (\alpha^{2} S_{2}) \frac{sh\lambda(h+y)}{sh\lambda h}] \cos S + 2\alpha\lambda\sigma_{0} a^{2} [C_{3} \frac{ch2\lambda(h+y)}{ch2\lambda h} + S_{3} \frac{sh2\lambda(h+y)}{ch2\lambda h} + S_{3} \frac{sh2\lambda(h+y)}{ch2\lambda h} , \qquad (19)$$

$$- (C_{1} + \frac{D + ch^{2}\lambda h}{D^{2}sh2\lambda h}) \frac{1}{sh^{2}\lambda h}] \sin 2S + \lambda\sigma_{0} a^{2} (\frac{3}{4} \frac{ch2\lambda(h+y)}{sh^{4}\lambda h} - \frac{1}{2} \frac{1}{sh^{2}\lambda h}) \cos 2S + [M_{2,0,0} + \frac{1}{2} \sigma_{0}\lambda a^{2} \frac{ch2\lambda(h+y)}{sh^{2}\lambda h}]$$

$$v_{t} = v_{e} + \left(\int^{t} u_{e} dt\right) \frac{\partial v_{e}}{\partial x} + \left(\int^{t} v_{e} dt\right) \frac{\partial v_{e}}{\partial y} = \sigma_{0} a \left[\left(1 - \alpha^{2} C_{2}\right) \frac{sh\lambda(h+y)}{sh\lambda h} - \left(\alpha^{2} S_{2}\right) \frac{ch\lambda(h+y)}{sh\lambda h}\right] \sin S + \sigma_{0} a \left[-\left(\alpha C_{1} + \alpha^{3} C_{3}\right) \frac{sh\lambda(h+y)}{sh\lambda h} - \left(\alpha^{2} S_{1} + \alpha^{3} S_{3}\right) \frac{ch\lambda(h+y)}{sh\lambda h}\right] \cos S + \frac{3}{4} \lambda \sigma_{0} a^{2} \frac{sh2\lambda(h+y)}{sh^{4}\lambda h} \times \sin 2S - 2\alpha\lambda\sigma_{0} a^{2} \left[C_{3} \frac{sh2\lambda(h+y)}{ch2\lambda h} + S_{3} \frac{ch2\lambda(h+y)}{ch2\lambda h} + S_{3} \frac{ch2\lambda(h+y)}{ch2\lambda h}\right], \quad (20)$$
$$- \left(\frac{\lambda(h+y)}{2Dsh2\lambda h} - \frac{1}{2}\right) \frac{1}{sh^{2}\lambda h} \left[\cos 2S\right]$$

Then, the displacement components X and Y of the fluid particle with initial average position (x, y) are obtained by direct integration as

$$\begin{split} X &= \int^{t} u_{l} dt \\ &= a\{[(-1+\alpha^{2}C_{2})\sin S + (\alpha C_{1}+\alpha^{3}C_{3})\cos S]\frac{\mathrm{ch}\lambda(h+y)}{\mathrm{sh}\lambda h} \\ &+ [\alpha^{2}S_{2}\sin S + (\alpha S_{1}+\alpha^{3}S_{3})\cos S]\frac{\mathrm{sh}\lambda(h+y)}{\mathrm{sh}\lambda h}\} \\ &- \frac{3}{8}\lambda a^{2}\frac{\mathrm{ch}2\lambda(h+y)}{\mathrm{sh}^{4}\lambda h}\sin 2S + \frac{1}{4}\lambda a^{2}\frac{1}{\mathrm{sh}^{2}\lambda h}\sin 2S \\ &+ [M_{2,0,0} + \frac{1}{2}\sigma_{0}\lambda a^{2}\frac{\mathrm{ch}2\lambda(h+y)}{\mathrm{sh}^{2}\lambda h}]t + \alpha\lambda a^{2}[C_{3}\frac{\mathrm{ch}2\lambda(h+y)}{\mathrm{ch}2\lambda h} \\ &+ S_{3}\frac{\mathrm{sh}2\lambda(h+y)}{\mathrm{ch}2\lambda h} - (C_{1} + \frac{D + \mathrm{ch}^{2}\lambda h}{D^{2}\mathrm{sh}2\lambda h})\frac{1}{\mathrm{sh}^{2}\lambda h}]\cos 2S \end{split}$$

$$Y = \int^{t} v_{l} dt$$

$$= a\{[-\alpha^{2}S_{2}\cos S + (\alpha S_{1} + \alpha^{3}S_{3})\sin S]\frac{\mathrm{ch}\lambda(h+y)}{\mathrm{sh}\lambda h}$$

$$+[(1-\alpha^{2}C_{2})\cos S + (\alpha C_{1} + \alpha^{3}C_{3})\sin S]\frac{\mathrm{sh}\lambda(h+y)}{\mathrm{sh}\lambda h}\}, \quad (22)$$

$$+\frac{3}{8}\lambda a^{2}\frac{\mathrm{sh}2\lambda(h+y)}{\mathrm{sh}^{4}\lambda h}\cos 2S + \alpha\lambda a^{2}[C_{3}\frac{\mathrm{sh}2\lambda(h+y)}{\mathrm{ch}2\lambda h}]$$

$$+S_{3}\frac{\mathrm{ch}2\lambda(h+y)}{\mathrm{ch}2\lambda h} - (\frac{\lambda(h+y)}{2D\mathrm{sh}2\lambda h} - \frac{1}{2})\frac{1}{\mathrm{sh}^{2}\lambda h}]\sin 2S,$$

where

....

$$C_{1} = \frac{\lambda^{2}(h+y)^{2}}{D \operatorname{sh} 2\lambda h} - \lambda(h+y) + \frac{1}{D^{2} \operatorname{th} \lambda h};$$
  

$$S_{1} = \frac{2(D+\operatorname{ch}^{2}\lambda h)}{D^{2} \operatorname{sh} 2\lambda h} \lambda(h+y) - 1$$

$$C_{2} = \frac{\lambda^{4}(h+y)^{4}}{2D^{2} \operatorname{sh}^{2} 2\lambda h} - \frac{\lambda^{3}(h+y)^{3}}{D \operatorname{sh} 2\lambda h} + (\frac{1}{2} + \frac{4D^{2} + 10D \operatorname{ch}^{2} \lambda h - 2D \operatorname{sh}^{2} 2\lambda h + 10 \operatorname{ch}^{4} \lambda h}{D^{4} \operatorname{sh}^{2} 2\lambda h}) \times \lambda^{2}(h+y)^{2} + \frac{4D + 6 \operatorname{ch}^{2} \lambda h}{D^{2} \operatorname{sh} 2\lambda h} \lambda(h+y) - 1 + \frac{k_{0,2}}{\lambda}$$

$$S_{2} = \frac{10(D + ch^{2} \lambda h)}{3D^{3} sh^{2} 2\lambda h} \lambda^{3} (h + y)^{3}$$
$$- \frac{5D + 2 ch^{2} \lambda h}{D^{2} sh 2\lambda h} \lambda^{2} (h + y)^{2}$$
$$+ (2 + \frac{5I^{2} - 4D}{D^{4}} - \frac{k_{0,2}}{\lambda})\lambda(h + y) - \frac{I}{D^{2}}$$

$$C_{3} = \frac{3(I^{4} - 1)}{4Dsh2\lambda h}\lambda^{2}(h + y)^{2} - \frac{3(I^{2} + 1)}{4sh^{2}\lambda h}\lambda(h + y)$$
  
+  $\frac{1}{16D^{2}}(I^{2} + 1) \times [12D^{2}I^{6} + (6D^{2} + 3D)I^{5}$   
-  $(48D^{2} - 24D)I^{4} - (18D^{2} + 2D - 22)I^{3}$   
+  $(36D^{2} - 24D)I^{2} - (D + 10)I]$   
+  $\frac{3[2D(I^{4} - 1) + (D + 1)(I^{3} + I)]}{4Dsh^{2}\lambda h}$ 

$$S_3 = \frac{3[2D(I^4 - 1) + (D + 1)(I^3 + I)]}{4D^2 sh^2 \lambda h} + \frac{3(I^4 - 1)}{4Dsh2\lambda h}$$

Substituting y = 0 into Equations (21) and (22), the displacement of a water particle on the free surface  $\xi$  (x-direction) and  $\zeta$  (y-direction) can be given as

$$\xi = x + (X)_{y=0}$$

$$= x + a\{[(-1 + \alpha^2 C_{20})\sin S + (\alpha C_{10} + \alpha^3 C_{30})\cos S]I$$

$$+ [\alpha^2 S_{20}\sin S + (\alpha S_{10} + \alpha^3 S_{30})\cos S]\}$$

$$- \frac{3}{8}\lambda a^2 \frac{ch2\lambda h}{sh^4\lambda h} \times \sin 2S + \frac{1}{4}\lambda a^2 \frac{1}{sh^2\lambda h}\sin 2S + [M_{2,0,0} \qquad (23)$$

$$+ \frac{1}{2}\sigma_0\lambda a^2 \frac{ch2\lambda h}{sh^2\lambda h}]t + \alpha\lambda a^2[C_{30} + S_{30}th2\lambda h$$

$$- (C_{10} + \frac{D + ch^2\lambda h}{D^2sh2\lambda h})\frac{1}{sh^2\lambda h}]\cos 2S,$$

$$\zeta = (Y)$$

$$= a\{[-\alpha^{2}S_{20}\cos S + (\alpha S_{10} + \alpha^{3}S_{30})\sin S]I + [(1 - \alpha^{2}C_{20})\cos S + (\alpha C_{10} + \alpha^{3}C_{30})\sin S]\}, \qquad (24)$$
$$+ \frac{3}{8}\lambda a^{2}\frac{sh2\lambda h}{sh^{4}\lambda h}\cos 2S + \alpha\lambda a^{2}[C_{30}th2\lambda h + S_{30} - (\frac{\lambda h}{2Dsh2\lambda h} - \frac{1}{2})\frac{1}{sh^{2}\lambda h}]\sin 2S,$$

where  $(\xi, \zeta)$  is the position of the fluid particle on the free surface (y = 0).

In the above expressions,  $C_{10}$  ,  $S_{10}$  ,  $C_{20}$  ,  $S_{20}$  ,  $C_{30}$  , and  $S_{30}$  are deduced from  $C_1$  ,  $S_1$  ,  $C_2$  ,  $S_2$  ,  $C_3$  , and  $S_3$ given in Equations (23) and (24), with the substitution y = 0, where the subscript 0 represents the free surface.

#### 2. Breaker characteristics

The increase in wave height owing to the shoaling effect becomes depth-limited as the wave propagates into shallow waters. The celerity is reduced; the particle velocity of the wave crest is faster than the wave celerity, and the wave breaks. To describe the breaking-wave mechanism, the kinematic stability parameter (K.S.P.) is applied, and the breaking-wave

criterion is

$$K.S.P. = \frac{u_{0b}}{C} = 1, \qquad (25)$$

where *C* is wave celerity and  $u_{0b}$  is the horizontal velocity of the particle at the wave crest.

In accordance with Equation (18), the breaking amplitude parameters  $a_b$  may be written as

$$a_{b} = \frac{a_{0}}{(D_{b} \tanh k_{b} h_{b})^{\frac{1}{2}}}$$
(26)

where  $D_b = (1 + \frac{2k_b h_b}{\sinh 2k_b h_b})$ , and the subscript *b* denotes the

breaking condition.

Up to the breaking point, the linear dispersion relation is still valid, and hence

$$\sigma^2 = gk_b \tanh k_b h_b; \ \sigma = k_b C_b; \ L_b = \frac{2\pi}{k_b}.$$
(27)

where  $L_b$  is the wavelength,  $k_b$  is the wave number,  $C_b$  is the celerity of the breaking wave, and  $h_b$  is the breaking-wave depth, respectively.

The Lagrangian horizontal velocities  $u_l$  evaluated at y = 0 can be simplified as

$$(u_{l})_{0} = \sigma_{0}a[(\alpha C_{10} + \alpha^{3}C_{30})I + (\alpha S_{10} + \alpha^{3}S_{30})]\sin S + \sigma_{0}a[(1 - \alpha^{2}C_{20})I - (\alpha^{2}S_{20})]\cos S + 2\alpha\lambda\sigma_{0}a^{2}[C_{30} + S_{30}th2\lambda h - (C_{10} + \frac{D + ch^{2}\lambda h}{D^{2}sh2\lambda h})$$
(28)  
$$\times \frac{1}{sh^{2}\lambda h}]\sin 2S + \lambda\sigma_{0}a^{2}(\frac{3}{4}\frac{ch2\lambda h}{sh^{4}\lambda h} - \frac{1}{2}\frac{1}{sh^{2}\lambda h})\cos 2S + [M_{2,0,0} + \frac{1}{2}\sigma_{0}\lambda a^{2}\frac{ch2\lambda h}{sh^{2}\lambda h}],$$

The breaker condition is defined by taking the extreme value of the horizontal velocity

$$\frac{\partial (u_l)_0}{\partial t} = 0 \tag{29}$$

The phase angle of the breaking wave  $S_b$  can be solved, and this condition also implies that the water surface elevation  $\zeta$  has an extreme value.

The resulting Lagrangian horizontal velocity of the breaking wave on the free surface, based on Eq. (28), is denoted as  $(u_l)_{0b}$ :

$$(u_{l})_{0b} = \sigma_{0}a_{b}[(\alpha C_{10b} + \alpha^{3}C_{30b}) \coth k_{b}h_{b}$$

$$+ (\alpha S_{10b} + \alpha^{3}S_{30b})]\sin S_{b} + \sigma_{0}a_{b}[(1 - \alpha^{2}C_{20b}) \coth k_{b}h_{b}$$

$$- (\alpha^{2}S_{20b})]\cos S_{b} + 2\alpha k_{b}\sigma_{0}a_{b}^{2}[C_{30b} + S_{30b}th2k_{b}h_{b}$$

$$- (C_{10b} + \frac{D_{b} + ch^{2}k_{b}h_{b}}{D_{b}^{2}sh2k_{b}h_{b}})\frac{1}{sh^{2}k_{b}h_{b}}]\sin 2S_{b}$$
(30)
$$+ k_{b}\sigma_{0}a_{b}^{2} \times (\frac{3}{4}\frac{ch2k_{b}h_{b}}{sh^{4}k_{b}h_{b}} - \frac{1}{2}\frac{1}{sh^{2}k_{b}h_{b}})\cos 2S_{b}$$

$$+ [\frac{-\sigma_{0}a_{b}^{2}}{2thk_{b}h_{b}} + \frac{1}{2}\sigma_{0}k_{b}a_{b}^{2}\frac{ch2k_{b}h_{b}}{sh^{2}k_{b}h_{b}}].$$

 $C_{10b}$ ,  $C_{20b}$ ,  $C_{30b}$ ,  $S_{10b}$ ,  $S_{20b}$ ,  $S_{30b}$ , and  $k_b$  are deduced from  $C_{10}$ ,  $C_{20}$ ,  $C_{30}$ ,  $S_{10}$ ,  $S_{20}$ ,  $S_{30}$ , and k, with the substitution  $a = a_b$ ,  $D = D_b$ ,  $k = k_b$ , and  $h = h_b$ , where the subscript *b* represents the breaking wave.

Thus, the breaking-wave criterion can be rewritten in the more specific form

$$K.S.P. = \frac{(u_l)_{0b}}{C_b} = 1.$$
 (31)

This criterion, after substituting the previous Lagrangian velocity, becomes

$$k_{b}a_{b}[(\alpha C_{10b} + \alpha^{3}C_{30b}) \coth k_{b}h_{b} + (\alpha S_{10b} + \alpha^{3}S_{30b})]\sin S_{b}$$

$$+k_{b}a_{b}[(1 - \alpha^{2}C_{20b}) \coth k_{b}h_{b}$$

$$-(\alpha^{2}S_{20b})]\cos S_{b} + 2\alpha \times k_{b}^{2}a_{b}^{2}[C_{30b} + S_{30b}th2k_{b}h_{b}$$

$$-(C_{10b} + \frac{D_{b} + ch^{2}k_{b}h_{b}}{D_{b}^{2}sh2k_{b}h_{b}}) \times \frac{1}{sh^{2}k_{b}h_{b}}]\sin 2S_{b}$$

$$+k_{b}^{2}a_{b}^{2}(\frac{3}{4}\frac{ch2k_{b}h_{b}}{sh^{4}k_{b}h_{b}} - \frac{1}{2}\frac{1}{sh^{2}k_{b}h_{b}}) \times \cos 2S_{b}$$

$$+[\frac{-k_{b}a_{b}^{2}}{2thk_{b}h_{b}} + \frac{1}{2}k_{b}^{2}a_{b}^{2}\frac{ch2k_{b}h_{b}}{sh^{2}k_{b}h_{b}}] = 1.$$
(32)

By utilizing Equations (27), (29), and (32), we can solve  $h_b$ ,  $k_b$ , and  $S_b$ . The maximum surface elevation  $\varsigma_{max}$ , which appears when a wave breaks, is given by

$$\begin{aligned}
\varsigma_{\max} &= a_b \{ [(1 - \alpha^2 C_{20b} - \alpha^2 S_{20b} \coth k_b h_b) \cos S_b \\
&+ [(\alpha C_{10b} + \alpha^3 C_{30b}) + (\alpha S_{10b} + \alpha^3 S_{30b}) \\
&\times \coth k_b h_b ] \sin S_b \} + \frac{3}{8} k_b a_b^{-2} \times \frac{sh2k_b h_b}{sh^4 k_b h_b} \cos 2S_b \\
&+ \alpha k_b a_b^{-2} [C_{30b} th2k_b h_b + S_{30b} \\
&- (\frac{k_b h_b}{2D_b sh2k_b h_b} - \frac{1}{2}) \frac{1}{sh^2 k_b h_b} ] \sin 2S_b.
\end{aligned}$$
(33)



Fig. 3. Relationship between  $H_{\rm b}/h_{\rm b}$  and  $h_{\rm b}/L_0$ 



Fig. 5. Relationship between  $H_b/H_0$  and  $H_0/L_0$  for  $\alpha = 1/50$ .

Similarly, we can solve the minimum elevation of breaking-wave depth  $\zeta_{\min}$ , and the breaking-wave height  $H_b$  is then derived as

$$H_{b} = (\varsigma_{\max} - \varsigma_{\min}) = (\varsigma_{b} - \varsigma_{\min})$$
(34)

where  $\zeta_b = \zeta_{\max}$ .

# V. THEORY VALIDATION AND DISCUSSION

Many factors influence wave breaking, including wave



Fig. 4. Relationship between  $H_b/H_0$  and  $H_0/L_0$  for  $\alpha = 1/20$ .



Fig. 6. Relationship between  $h_{b}/H_{0}$  and  $H_{0}/L_{0}$  for  $\alpha$  = 1/20.

steepness, wave height, and bottom slope. The breaking phenomenon is so complicated that most related studies have described breaking-wave characteristics only from experimental studies and empirical or semiempirical formulae calibrated from laboratory data. However, this paper presents a theoretical investigation of nonlinear surface-wave propagation over a sloping bottom and compares the results of the present theory with previously published results (Fig. 3 to Fig. 8).

Fig. 3 to Fig. 5 demonstrate that  $H_b/h_b$  decreases as  $h_b/L_0$  increases and that  $H_b/H_0$  decreases as  $H_0/L_0$  increases. Fig.

3 shows that the present analytical solutions follow the same trend that other research results do. Fig. 4 to Fig. 8 also demonstrate that the theoretical modeling results follow the same trend that experimental measurements in other studies have.

# **VI. CONCLUSIONS**

This paper provides an analytical solution for breaker characteristics. The solution has not been appropriately addressed in other studies. The present theoretical solutions are compared with the results of other scholars, as shown in Fig. 3 to Fig. 5. They also show that the present analytical solutions follow the same trend that previously published results have. The main outcomes of this study are as follows:

- This paper provides a theoretical solution for a wave propagating over a uniformly sloping bottom, from deep water to shallow water, until the wave breaks. The theoretical results of this study reveal characteristics of breaking waves that have not been described in other studies.
- 2. The respective theoretical modeling requires only knowledge of the incident wave conditions and of the bottom slope for breaking-wave characteristics to be obtained.

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